Tsunami identification from surface data: an illustration of the Tikhonov-Morozov strategy

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A simple ocean model (time domain)

Velocity potential $\varphi(x, z, t)$, free surface perturbation s(x, t), vertical bottom displacement $\zeta(x, t)$



Potential φ satisfies:

$$\begin{cases} (1/c^2)\partial_t^2 \varphi - \Delta \varphi &= 0 & \text{in } \Omega \times (0, +\infty) \\ \partial_t \varphi + g \, s &= 0 & \text{on } \Gamma_0 \times (0, +\infty) \\ \partial_z \varphi - \partial_t s &= 0 & \text{on } \Gamma_0 \times (0, +\infty) \\ \partial_z \varphi &= \partial_t \zeta & \text{on } \Gamma_{-H} \times (0, +\infty) \end{cases}$$

The frequency domain: $\varphi(x, z, t) = u(x, z)e^{-i\omega t}$

Velocity potential u(x, z), free surface perturbation $\eta(x)$, vertical bottom displacement $\chi(x)$



Potential u satisfies (T_{\pm} : Dirichlet-to-Neumann operators):

$$\begin{aligned} \Delta u + (\omega^2/c^2)u &= 0 & \text{in } \Omega^R \\ \partial_z u - (\omega^2/g)u &= 0 & \text{on } \Gamma_0^R \\ \partial_z u &= -i\omega\chi & \text{on } \Gamma_{-H}^R \\ \pm \partial_x u - T_{\pm}u &= 0 & \text{on } \Sigma_{\pm R} \end{aligned}$$

The tsunami inverse problem

From the measurement of the **free surface** $\eta = (i\omega/g)u$ on Γ_0 , find the **tsunami** χ on $\Gamma_{-H} \longrightarrow$ equivalent to a **strongly ill-posed** Cauchy problem for the Helmholtz equation

Compute u such that:

$$\begin{aligned}
\Delta u + (\omega^2/c^2)u &= 0 & \text{in } \Omega^R \\
\partial_z u - (\omega^2/g)u &= 0 & \text{on } \Gamma_0^R \\
u &= u^{\text{mes}} & \text{on } \Gamma_0^R \\
\pm \partial_x u - T_{\pm}u &= 0 & \text{on } \Sigma_{\pm R}
\end{aligned}$$

with $u^{\text{mes}} := (g/i\omega)\eta^{\text{mes}}$

then compute $\chi = (i/\omega)\partial_z u$ on Γ^R_{-H}

The tsunami inverse problem (cont.)

An abstract framework

Spaces:

$$V = H^{1}(\Omega^{R}), \quad M = \{ \mu \in H^{1}(\Omega^{R}), \ \mu|_{\Gamma^{R}_{-H}} = 0 \}, \quad O = L^{2}(\Gamma^{R}_{0})$$

Operators: $B: V \to M$ defined by

$$b(v,\mu) = \int_{\Omega^R} \left(\nabla v \cdot \nabla \overline{\mu} - (\omega^2/c^2) v \overline{\mu} \right) dx dz - (\omega^2/g) \int_{\Gamma_0^R} v \,\overline{\mu} \, dx$$

$$- \langle T_+ v, \overline{\mu} \rangle_{H^{-1/2}(\Sigma_R), \tilde{H}^{1/2}(\Sigma_R)} - \langle T_- v, \overline{\mu} \rangle_{H^{-1/2}(\Sigma_{-R}), \tilde{H}^{1/2}(\Sigma_{-R})}$$

$$= (Bv, \mu)_M$$

 $C: V \to O$ is the trace operator,

 $A: V \to M \times O$ is defined as A = (B, C)

Data: is given by $F = (0, u^{\text{mes}}) \in M \times O$

The tsunami inverse problem

The inverse problem: for $F \in M \times O$, find $u \in V$ such that Au = F (A is injective, dense range but not onto) \longrightarrow we apply the Tikhonov-Morozov strategy for A

Mixed Tikhonov formulation: find $(u_{\varepsilon}, \lambda_{\varepsilon}) \in V \times M$ such that

$$\begin{cases} \varepsilon(u_{\varepsilon}, v)_{H^{1}(\Omega^{R})} + \int_{\Omega^{R}} \nabla \lambda_{\varepsilon} \cdot \nabla \overline{v} - (\omega^{2}/c^{2})\lambda_{\varepsilon}\overline{v} \, dx dz - (\omega^{2}/g) \int_{\Gamma_{0}^{R}} \lambda_{\varepsilon} \overline{v} \, dx \\ -\langle \overline{T_{+}}\lambda_{\varepsilon}, \overline{v} \rangle_{H^{-1/2}(\Sigma_{R}), \tilde{H}^{1/2}(\Sigma_{R})} - \langle \overline{T_{-}}\lambda_{\varepsilon}, \overline{v} \rangle_{H^{-1/2}(\Sigma_{-R}), \tilde{H}^{1/2}(\Sigma_{-R})} \\ + \int_{\Gamma_{0}^{R}} u_{\varepsilon} \overline{v} \, dx = \int_{\Gamma_{0}^{R}} u^{\text{mes}} \overline{v} \, dx, \qquad \forall v \in V \\ \int_{\Omega^{R}} \nabla u_{\varepsilon} \cdot \nabla \overline{\mu} - (\omega^{2}/c^{2}) u_{\varepsilon} \overline{\mu} \, dx dz - (\omega^{2}/g) \int_{\Gamma_{0}^{R}} u_{\varepsilon} \overline{\mu} \, dx \\ -\langle T_{+}u_{\varepsilon}, \overline{\mu} \rangle_{H^{-1/2}(\Sigma_{R}), \tilde{H}^{1/2}(\Sigma_{R})} - \langle T_{-}u_{\varepsilon}, \overline{\mu} \rangle_{H^{-1/2}(\Sigma_{-R}), \tilde{H}^{1/2}(\Sigma_{-R})} \\ - \int_{\Omega^{R}} \nabla \lambda_{\varepsilon} \cdot \nabla \overline{\mu} \, dx dz = 0, \qquad \forall \mu \in M \end{cases}$$

Numerical experiments

Artificial data obtained with the complete model

Inverse problem solved with the gravity model (case $c \to +\infty$) or the acoustic model (case $g \to 0$).



Exact data, 1% noise, 5% noise, 10% noise

Gravity case for $\omega = 3$, comparison between retrieved χ (dashed line) and exact χ (continuous line) for various amplitudes of noise

Numerical experiments (cont.)

Artificial data obtained with the complete model

Inverse problem solved with the gravity model (case $c \to +\infty$) or the acoustic model (case $g \to 0$).



Exact data, 1% noise, 5% noise, 10% noise

Acoustic case for $\omega = 20$, comparison between retrieved χ (dashed line) and exact χ (continuous line) for various amplitudes of noise

Numerical experiments (cont.)

How good is the Morozov choice ?



Error $\|\chi_{\varepsilon}^{\delta} - \chi\|_{L^{2}(\Gamma_{0}^{R})}$ in the acoustic case as a function of ε **Blue**: no noise, **Red**: 5% noise.

The vertical dashed thin/thick lines corresponds to the Morozov choice following deterministic/probabilistic procedure to lift Neumann data